

Planning for gypsum geohazards in Lithuania and England

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Abstract

The rapid underground dissolution of gypsum, and the evolution of the gypsum karst in Lithuania and England, results in subsidence problems which can make construction difficult. The natural dissolution yields sulphate-rich groundwater of poor quality and the karst is susceptible to the rapid transmission of pollutants.

In the north of Lithuania gypsum karst is developed in Devonian gypsum. Here the towns of Birai, Pasvalys and the surrounding countryside suffer subsidence and some buildings have been damaged. The majority of the potable water in these areas is derived from groundwater extracted from sandstone sequences that underlie the gypsum. In Lithuania conservation measures have been introduced to control agriculture and prevent pollution of the gypsum karst. These measures include environmentally-friendly farming, restrictions on land use and exclusion zones around subsidence hollows.

In England subsidence caused by the dissolution of Permian gypsum has caused severe problems in the vicinity of the town of Ripon. Numerous buildings have been damaged and new sites are difficult to develop. Here formal planning regulations have recently been introduced to help to protect against the worst effects of subsidence resulting from gypsum dissolution. © 1999 Elsevier Science B.V. All rights reserved.

Keywords: Gypsum karst; Subsidence; Hazard; Planning

1. Introduction

Gypsum is a very soluble mineral which can dissolve at a rapid rate. Where natural dissolution of exposures has occurred adjacent to rivers it is common for one metre of gypsum to be dissolved away in ca 1 year (James et al., 1981; James, 1992). Where this dissolution has occurred or is occurring, underground cave systems can develop such as those explored in the Ukraine

(Andrajchouk and Klimchouk, 1993), Germany (Biese, 1931; Pfeiffer and Hahn, 1972) and Spain (Pulido-Bosch and Calaforra, 1993). Because the dissolution rate is so rapid gypsum cave systems can enlarge at a considerable rate, ultimately become unstable, and collapse causing subsidence problems at the surface. The mechanism of collapse causes sub-vertical breccia pipes to develop with subsidence hollows where these break through to the surface (Cooper, 1986, 1988, 1995).

In most countries where gypsum occurs in contact with water there are associated subsidence problems. In farmland these are inconvenient, but

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in urban areas they constitute a geological hazard that can seriously affect development and human safety. Gypsum geohazards affect the towns of Biržai and Pasvalys in Lithuania (Paukštys, 1996) and Ripon and Darlington in England (Cooper, 1995; and references therein). Elsewhere in Europe gypsum geohazards are present in many towns and cities. For example, in Spain they have been recorded in the city of Zaragoza (Benito et al., 1995) and the town of Calatayud (Gutiérrez, 1996); in France they affect the outskirts of Paris (Toulemont, 1984), and in Stuttgart, Germany and many towns peripheral to the Hartz Mountains suffer subsidence problems (Pfeiffer and Hahn, 1972; Ströbel, 1973). In addition to these examples, gypsum dissolution and subsidence affects many more urban and rural areas in these and many other countries. Some of these areas may be the sites of future roads, reservoirs or urban growth. Thus, an appreciation of gypsum geohazards is important for planning and development on a national, provincial and local scale.

In addition to the problems of subsidence, some countries such as Lithuania rely heavily on groundwater for their potable water supplies (Klimas and Paukštys, 1993; Paukštys, 1996). Despite its mineral content, sulphate-rich water associated with the gypsum karst areas is sometimes the only water supply that can be obtained. Extraction of this water, or water from aquifers in contact with the gypsum karst, can result in subsidence both from drawdown of the water table and from the increased dissolution of gypsum, especially in the vicinities of boreholes. Drawdown of the water table causes a loss of hydrostatic buoyancy and the effective weight of cavity fill increases which can cause collapse. Drawdown can also wash material deeper into cavities and aggravate the subsidence problems. Another problem is that the rapid passage of groundwater, both through limestone and gypsum karst, can lead to the swift transmission of pollutants from their source to a potable water supply (Klimas and Paukštys, 1993). Gypsum karst, therefore, demands careful management and possible protection if the land and water associated with it are to be used to their full potential.

2. Geology, subsidence and hydrological considerations in the gypsum karst of Lithuania

2.1. Geology

The karst area of northern Lithuania covers ca 20 000 km², of which ca 1000 km² are composed of gypsum karst. The gypsum karst area is well developed around the towns of Pasvalys and Biržai extending northwards into Latvia (Fig. 1).

The gypsum in northern Lithuania is of late Devonian age. It occurs in the Tatula Formation where two main gypsiferous sequences are interbedded with dolomites and marls (Table 1); gypsum composes some 70% of the sequence. The gypsum sequence is underlain by dolomites of the Pliavinias Formation including the thin, low-permeability Jara member which overlies the sandstone aquifer of the Šventoji Formation and the argillaceous sandstones aquifer of the Upninkai Formation. Below this the Narva Clay Formation forms a regional aquiclude that limits the karstic aquifer basin. The gypsum karst is locally covered by the thin (3–9 m) dolomite of the Ystra Formation, but over most of the outcrop it is mantled only by Quaternary deposits which are up to 20 m in thickness. The Quaternary deposits are glacial tills with lenticular belts of sand which form minor, local near-surface aquifers.

2.2. Subsidence

The gypsum karst area of northern Lithuania has well-developed sinkholes that range in density from 20 to 200 km⁻². They range in size from a few metres to 60 m in diameter and are up to 12 m in depth (Marcinkevičius and Bucevičiute, 1986). The sinkholes are concentrated in areas related to the valleys and water divides of the Muša, Levuo and Pyvesa rivers, with the greatest density of sinkholes on the Kirkilai geological reserve. Here, in the bottom of one sinkhole, there is also a small accessible cave in the gypsum. This cave, dedicated as a geological monument, is up to 3.1 m high, with 46 m of accessible passages. It is of phreatic origin with water-eroded scallops on the roof, but it is now only half full of water (Laiconas, 1979). Elsewhere in this area, stream sinks of moderate

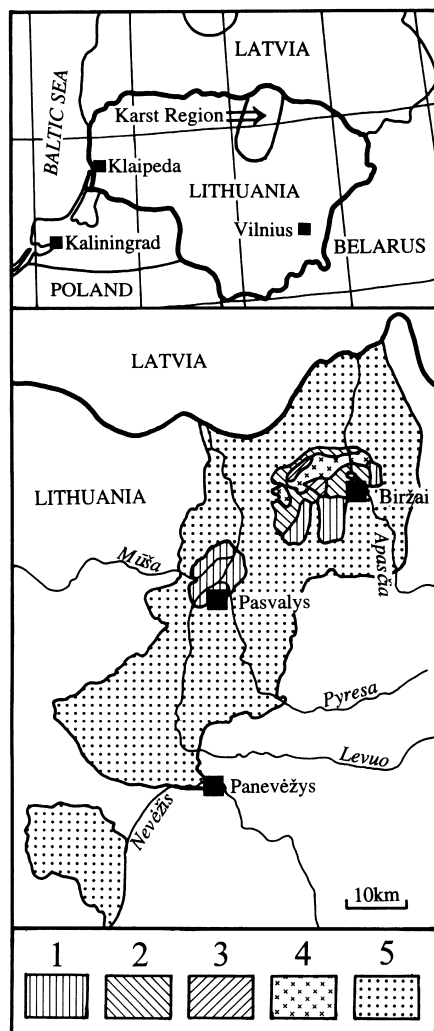


Fig. 1. The location of the gypsum karst area in northern Lithuania and the agricultural protection zonation. 1, Land group 1, <20 sinkholes/100 ha; 2, land group 2, 20–50 sinkholes/100 ha; 3, land group 3, 50–80 sinkholes/100 ha; 4, land group 4, >80 sinkholes/100 ha; 5, karst protection zone (100 ha = 1 km²).

size, such as the sinking of the 8 km long Požemis stream, indicate more extensive cave development. Prolific springs, common in the gypsum karst area, also suggest cave development. Major karst springs occur along the Levuo River at Pasvalys town, the Orija river near Berklainiai Village and the Apašėia river near Draseikiai Village.

The majority of the sinkholes (61%) are oval in

shape and their long axes correlate to the main joint directions in the gypsum, which are to the northeast, north and east (Marcinkevičius and Bucevičiute, 1986; Bucevičiute and Marcinkevičius, 1992). Lines of hollows also appear to relate to these joint directions. In the area of active karst >8500 sinkholes are present in an area of 400 km² (Bucevičiute and Marcinkevičius, 1992). By analogy with gypsum karst elsewhere it may be expected that many of the subsidence features are underlain by breccia pipes that extend to the base of the gypsum. This is corroborated by boreholes in the gypsum karst, which penetrate cavities, collapsed strata and washed-in materials within the gypsum. The deepest breccia pipe so far found, in the Radviliškis region, extended to 96.4 m (Bucevičiute and Marcinkevičius, 1992).

Active gypsum dissolution is indicated by the high concentrations of sulphates in groundwater coming from the various karst springs and also in the surface water of the main drainage courses such as the River Tatula. Active gypsum dissolution is also shown by the continuing collapse of the gypsum karst and the development of sinkholes, although their development may be aggravated by water extraction and changes in the water table levels. Where this collapse has occurred in the urban areas of both Pasvalys and Biržai, damage has ensued.

2.3. Hydrogeological considerations

In the gypsum karst area of northern Lithuania the main water supply comes from the Devonian aquifers. They are exploited through ca 600 bored wells with individual yields of between 10 and 50 m³ day⁻¹. In addition to the scattered wells, boreholes at the waterworks of Biržai and Pasvalys pump 2000 and 2600 m³ day⁻¹ from the aquifers. These large-scale withdrawals have caused a draw-down in the water table of 7.5 m since 1970 at Pasvalys, and 8 m since 1961 at Biržai. The karst area is currently monitored by GROTA, under the auspices of the Tatula Board (set up by Government decree), at 50 drilled wells, dug wells and karst springs. The detailed results of long-term monitoring and complex mathematical modelling of the karst aquifers are presented by

Table 1

The sequence of Devonian rocks in the north of Lithuania and their main lithological and hydrological properties

	Stage	Formation	Map code/ member	Thickness metres	Description	Hydrological properties	
Devonian	Frasnian	Ystra	Q	0–20	Glacial till and sands	Minor aquifer (especially sands); fresh water TDS 0.5–0.8 g l ^{−1} ; commonly polluted	
			D _{3ys}	3–9	Dolomite; fissured	Karst aquifer; very hard mineralised water TDS 1.5–2.4 g l ^{−1} sulphate and calcium-rich, commonly polluted	
			Tatula	D _{3tt} ^a	11–15	Gypsum intercalated with marl and dolomite	Aquifer; mainly fresh bicarbonate–calcium–magnesium water TDS 0.5–0.8 g l ^{−1} , areas with sulphate-rich water from above, some pollution
			D _{3tt} ^k	3–7	Marl		
			D _{3tt} ^p	13–24	Gypsum intercalated with marl and dolomite		
			Pliavinias	D _{3kp} /Kupiškis	6–12	Dolomite; fissured	
				D _{3ss} /Suosa	13–18	Dolomite; clayey	Aquitard
				D _{3j} /Jara	2.0–2.9	Dolomite and marl	
			Šventoji	D _{3šv}	90	Sandstone, coarse-grained intercalated with siltstone and fine-grained sandstone	
							Good aquifer; large amounts of fresh water, TDS 0.2–0.6 g l ^{−1} ; traces of sulphates and nitrogen show some local connection with the overlying aquifers
Givetian	Upninkai	D _{2up}	70–110	Sandy mudstone and mudstone			
Eifelian	Narva	D _{2nr}	~100	Calcareous mudstone	Low conductivity major regional aquiclude		

Paukštys (1996). In addition to the effects of local water extraction, up to 110 000 m³ day⁻¹ is forecast to be withdrawn by the large town of Panevezys, 40–60 km south of the gypsum karst area. This amount of pumping could have serious future drawdown effects on the gypsum karst water. Because drawdown is likely to aggravate the subsidence problems, future water development from the near-surface gypsum–dolomite karst aquifers has already been prohibited, but this remote withdrawal could also be dangerous. In addition to subsidence caused by drawdown, the lowering of the water table allows more aggressive groundwater to enter the gypsum karst. Approximate calculations carried out using the Lithuanian data show that the lowering of groundwater by 1 m increases the gypsum deficiency by

0.4 g l⁻¹ (Paukštys, 1996). In addition to the degradation of potable water, laboratory experiments and detailed chemical modelling show that common fertilizer compounds used in the gypsum karst area can lead to enhanced gypsum dissolution (Paukštys, 1996). The preferred aquifer for large-scale water development is the deeper Šventoji and Upninkai formations. These contain the best quality water, but with continued development are themselves becoming slightly contaminated with water drawn down from the overlying aquifers.

In order to categorize the susceptibility of the gypsum karst to pollution, 19 variables were analysed and classified; this is the grade method of Dublinskij et al. (1990). It involved defining the controlling parameters of the karst system. Solubility was defined by four factors: lithology;

thickness of gypsum deposits; content of soluble material; and geological structure. Permeability was defined by five factors: lithology of overlying sediments; thickness of overlying sediments; density of sinkholes; dip of the karst rocks and coefficient of transmissivity. The availability of groundwater was characterized by six factors: amount of effective precipitation; surface runoff; subsurface runoff; downward infiltration; seepage from neighbouring aquifers; and groundwater gradient. The aggressivity of the groundwater was characterized by three factors: water saturation degree as TDS; temperature; and pH. In addition to these factors, the number of old karst features (breccia pipes and debris-filled areas) was also considered. Using this technique, an integral grade scale was determined by summing up the separate active factors of the karst terrain. This allowed the karst to be divided into areas of weak karst (32–42 grades); medium karst (43–49 grades) and high karst (50–59 grades). These grades relate closely to the classification (Fig. 1) of the karst lands used for agricultural protection (Paukštys, 1996).

3. Geology, subsidence and hydrological considerations in the gypsum karst of England

3.1. Geology

In England gypsum karst and subsidence problems are mainly developed in the Permian sequence in northeast England (Fig. 2). Gypsum is present in the Edlington and Roxby formations from just north of Doncaster, through Ripon to Darlington and Hartlepool. Up to 40 m of gypsum are present in the Edlington Formation and 10 m in the Roxby Formation (Table 2). Both these gypsum sequences rest on dolomite aquifers and are capped by a marl sequence. However, in the subsidence-prone areas the amount of dissolution and collapse is so great that the marls are perforated by subsidence pipes and form very ineffective aquicludes. The Permian sequence is capped by the Triassic Sherwood Sandstone Group, a major regional aquifer. In addition to the Permian gypsum the majority of the English-mined gypsum is in the Triassic Mercia Mudstone Group (Fig. 2). Some subsidence has been noted associated with this

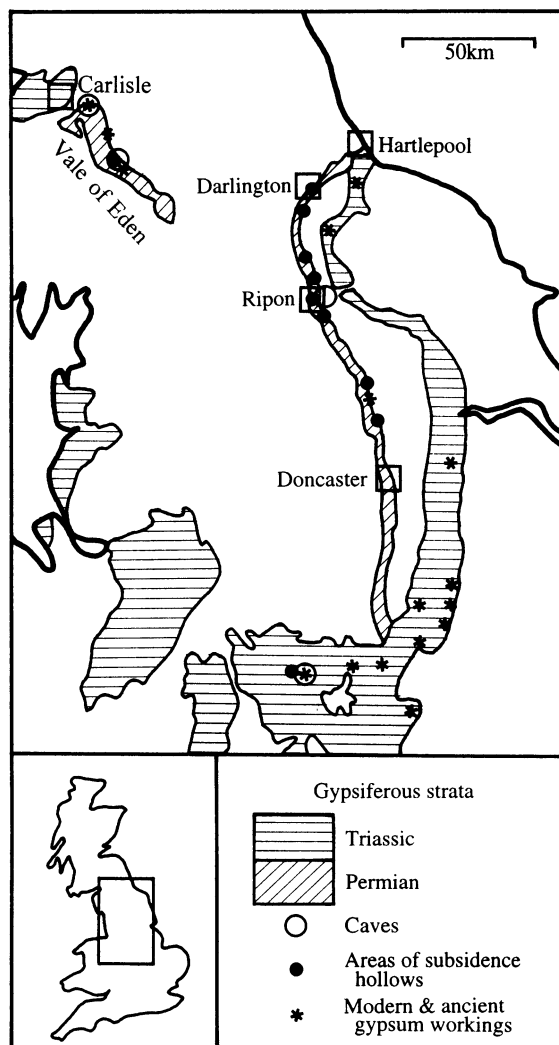


Fig. 2. The distribution of the main gypsiferous strata in England showing the location of past and present mines (*), caves (○) and subsidence hollow areas (●).

gypsum, but since the gypsum is sandwiched in mudstone aquicludes, subsidence is much more restricted in area. However, the dissolution of gypsum from the near-surface mudstones has considerably disrupted the fabric of the upper part of the mudstone sequence. This disruption and associated weathering have commonly resulted in the deposits presenting difficult ground for civil engineering purposes.

Table 2

The sequence of the Permian and Triassic rocks in northeastern England and their main lithological and hydrological properties

	Formation/group	Thickness (m)	Description	Hydrological properties
Triassic	Sherwood Sandstone Group	300	Red sandstone with subordinate mudstone beds, especially near base	Major regional aquifer TDS $0.15\text{--}0.3\text{ g l}^{-1}$ mainly as CaCO_3
Permian	Roxby Formation	Up to 26	Red-brown calcareous mudstone (marl) with up to 10 m of gypsum at base	Very leaky aquitard with a gypsum karst aquifer at base; sulphate-rich water
	Brotherton Formation	8–14	Calcitic dolomite, mainly in thin beds	Aquifer TDS $\sim 0.5\text{ g l}^{-1}$ as CaCO_3 ; sulphate-rich in places
	Edlington Formation	Up to 50	Red-brown calcareous mudstone (marl) with up to 30–40 m of gypsum at base	Very leaky aquitard with a gypsum karst aquifer at base TDS $0.8\text{--}2.0\text{ g l}^{-1}$ mainly as sulphate
	Cadeby Formation	Up to 65	Dolomitic limestone, commonly massive, but locally porous and leached	Major local aquifer TDS $0.2\text{--}0.5\text{ g l}^{-1}$ as CaCO_3

3.2. Subsidence

The two gypsum sequences of the Edlington and Roxby formations rest on the limestone aquifers of the Cadeby and Brotherton formations respectively. The limestone dip slopes act as catchment areas, and the water is fed down-dip into the gypsiferous sequences before escaping into a major buried valley along the line of the River Ure (Cooper and Burgess, 1993). Complex cave systems are developed in the gypsum, and artesian sulphate-rich springs are locally present. Because of the thickness of gypsum present, the caves are large, and surface collapses up to 30 m across and 20 m deep have been recorded. The subsidence is not random but occurs in a reticulate pattern related to the jointing in the underlying strata (Cooper, 1986, 1989). Around Ripon a significant subsidence occurs approximately every year (Cooper, 1995). The times of the subsidence events show that some zones of subsidence are more active than others. Furthermore, areas bounding the Ure valley are more subsidence-prone due to the localized escape of cave water into the buried valley gravels. In England, gypsum caves and subsidence are not confined to Ripon; the subsidence-prone belt is ca 3–4 km wide and extends from near Doncaster to Hartlepool. Several areas along this belt suffer gypsum-related subsidence,

though none is as severe as Ripon. Subsidence also affects the Darlington area, but in the urban district of this town, the problems are lessened by the presence of thick Quaternary glaciogenic deposits (Cooper, 1995).

3.3. Hydrogeological considerations

The natural dissolution of gypsum results in large quantities of sulphate in the groundwater. Consequently, in and around Ripon, many springs and waters in the glacial deposits are high ($0.8\text{--}2.0\text{ g l}^{-1}$) in sulphate. Calculations suggest that the volume of gypsum being dissolved naturally each year at Ripon is ca $120\text{ m}^3\text{ km}^{-2}$. However, north of Ripon a figure of ca $1000\text{ m}^3\text{ km}^{-2}$ may have been removed since the last (Devensian) ice age. In addition to this natural dissolution, extraction of groundwater high in sulphates can remove considerable volumes of gypsum from underground. It is estimated (Cooper, 1988) that the volume of gypsum removed by a group of boreholes in a subsidence damaged area of Ripon extracting $212\,000\text{ m}^3$ of water a year was ca $200\text{ m}^3\text{ year}^{-1}$. It is likely that much of the dissolution represented enlargement of joints over a considerable area. However, in the vicinity of the boreholes, where rapid groundwater flow occurs, severe dissolution of the gypsum beds

could have occurred. In addition to the dissolution problems, the resedimentation of Quaternary deposits into the gypsum karst may have resulted in surface subsidence in the former. This could also have been enhanced by localized lowering of the water table. Similar problems of glacial deposits being displaced into gypsum karst and causing subsidence have been suggested as a mechanism for the development of subsidence in the Darlington area (Cooper, 1995).

4. Planning in, and management of, gypsum karst areas

4.1. Hazard avoidance, the most cost-effective form of planning

The use of special building and development techniques, such as those outlined below, are expensive. If funding is not available, remediation or control schemes cannot be implemented, and it is impossible to legislate for special construction regulations. However, planning to avoid the worst of the hazardous areas and to limit the aggravation of the subsidence problems can be very cost-effective. The winners are people who avoid paying for constructions that subsequently fail, the losers are those with land that becomes less valuable for development. Avoidance of the worst areas causes less planning blight than developing and suffering severe subsidence and destruction of property and infrastructure.

It is largely impossible to avoid all development within the gypsum karst areas. In towns such as Ripon, where the margin of the subsidence belt runs through the town, it may be possible to encourage more development outside of the subsidence belt. However, within the subsidence areas, the first principle of avoiding gypsum geohazards is generally not to build in existing subsidence hollows. This is because they may still be unstable, they may have ongoing dissolution below them, or they may be filled with poorly consolidated deposits (or waste materials) with a low bearing strength. The second principle is not to build on the margins of the existing hollows, or between hollows in linear belts. This is because the collapse

of a hollow can lead to the choking of the underlying cave system. When this happens, the dissolution area can be pushed to the margins of the collapse and affect the adjacent ground. In this way subsidence hollows commonly occur in lines or close groupings. The third principle is to avoid the most active areas where the majority of the recent subsidence hollows have occurred.

4.2. Development and construction of buildings

The construction of buildings within gypsum karst requires special measures. In England the Government's Department of the Environment and Harrogate Borough Council (the local council to the Ripon area) have recently commissioned a report on planning and development in the subsidence-prone area (Thomson et al., 1996). The report approaches the problems on two fronts, construction and planning. For construction it reviews the problem and gives some possibilities for the types of foundations suitable for use in subsidence-prone areas. Options include raft foundations, jackable foundations and reinforced strip foundations. The report reiterates the difficulties and dangers of piling into gypsum karst, or of trying to improve the ground by grouting. These factors were discussed by Cooper (1995). Another approach to development is the use of extended foundations such as those suggested by Sorochan et al. (1985) or the construction of properties on linked foundations to prevent individual houses collapsing into subsidence hollows. In addition to these measures, precautions to protect services such as gas, water, electricity and sewerage, are also desirable. These precautions could include flexible pipe work, flexible connections and protection such as geogrid materials or reinforced supports.

The second approach to the subsidence problem recommended by Thomson et al. (1996) is through the planning regulatory process. To support this process the Ripon area has been divided into three development control zones:

- (A) no known gypsum present;
- (B) some gypsum present at depth;
- (C) gypsum present and susceptible to dissolution.

Within zone A no special planning constraints would be imposed. In zone B, where the risk of subsidence is small, a ground-stability report prepared by a competent person would usually be required, and the problem should be considered in local planning. The zone C area would be potentially subject to significant constraints on development, and local planning should take these into account. Also within this zone, development is subject to controls. A ground-stability report prepared by a competent professional person would normally be required before planning applications for new buildings, or change of use of buildings are determined. In most cases this report would need to be based on a geotechnical desk study and a site appraisal, followed by a program of ground investigation designed to provide information needed for detailed foundation design (unless this information, such as boreholes, exists from a previous study). Where planning consent is given, it may be conditional on the implementation of approved foundation or other mitigation measures, designed to minimize the impact of any future subsidence activity. One key to the implementation of this approach is the use of a proforma checklist to be completed and signed by a competent professional person. For the UK a competent person is defined in the report as a Geotechnical Specialist who is “a Chartered Engineer or Chartered Geologist, with a postgraduate qualification in geotechnical engineering or engineering geology, equivalent at least to an MSc, and with three years of post-Charter practice in geotechnics; or a Chartered Engineer or Chartered Geologist with five years of post-Charter practice in geotechnics”. In addition to these qualifications, it is also desirable that the practitioner has experience with the problems, although this is not formally stated. This procedure has been adopted by the Harrogate Borough Council, but it is likely to be subject to minor modifications as experience is gained.

4.3. Development and construction of roads and bridges

Sudden failure of roads over natural and man-made cavities has led to collapses in which vehicles have fallen into the resultant cavity. It is largely

impractical to engineer roads with design parameters of sufficient strength to span the larger subsidence features. Even if this could be done, the removal of support from beneath such structures could ultimately result in subsidence features migrating and the structures themselves failing catastrophically in a much larger way than non-protected structures. One practical approach that was adopted for a new bypass at Ripon was to incorporate several layers of geogrid material into the embankments of the road. If a subsidence develops beneath the road, the area of the subsidence will sink, but should not fail catastrophically. When subsidence occurs, its location will be obvious and some remedial measures can be undertaken. The use of geogrid materials is also a satisfactory method of protecting parking lots and public spaces.

The development of bridges in such situations is difficult. At Ripon the new road bridge has been built on the principle of having sacrificial supports. The deck of the bridge has been strengthened and built as a continuous structure so that the loss of support of any one upright will not cause it to collapse. A system of monitoring the loads on each support has been built into the bridge, and a warning system has been installed to warn of any pier failure. In addition to these measures, extending the foundations of each pier laterally to an amount which could span the normal-sized collapses would give an added degree of security.

4.4. Water extraction

Some details of the dangers of water drawdown and the active dissolution of gypsum are given above, in Cooper (1988) and in Paukštys (1996). Because it is possible to enhance both the local dissolution and cause of subsidence by the drawdown of the water table levels, careful monitoring and regulation of water extraction is essential in gypsum karst areas. In England and Lithuania attention is generally paid to water quality, but not so much thought is given to the subsidence implications of water extraction.

Another factor that must be considered is the effect of pollution on gypsum karst. Because the gypsum karst has rapid transmissivity in fissures

and caves, it is important to consider the implications of agriculture and waste disposal on water quality. If the water is used only for irrigation, a moderate content of nitrate and phosphate may not be immediately problematical. If the water is to be used as a potable supply, then rapid fluctuations in contaminants may occur. In such areas, careful consideration should also be given to protecting the gypsum karst from accidental contamination by spillage of chemicals, poor containment of farm wastes and foul water disposal.

4.5. Karst water protection and agriculture

In the Lithuanian karst area, 27 600 ha (276 km²) of intensive karst with strict agricultural limitations and 166 000 ha (1660 km²) of karst protection zone have been designated by government decree. Within this area (Fig. 1), four divisions of agricultural land use have been defined based mainly on the number of sinkholes per square kilometre. The categories and restrictions imposed are:

- (1) Land group 1 (up to 20 sinkholes/100 ha). Grain crops should compose at least 50% of arable lands, perennial grass 40% and root crops (potatoes and sugar beet) not more than 10%. Fertilizers are limited to a maximum of 90 kg ha⁻¹ of nitrogen/phosphorus/potassium (NPt active ingredients) and 80 t ha⁻¹ of manure. Triazinic herbicides and chlororganic insecticides are prohibited.
- (2) Land group 2 (20–50 sinkholes/100 ha). Grain crops should compose 43% of arable lands and perennial grass 57%. Root crops (potatoes and sugar beet) are prohibited as is the setting up of new orchards and gardens. Fertilizers are limited to a maximum of 60 kg ha⁻¹ of NPt and 60 t ha⁻¹ of manure.
- (3) Land group 3 (50–80 sinkholes/100 ha). Perennial grass and pastures only are allowed. Fertilizers are limited to a maximum of 60 kg ha⁻¹ NPK. Mineral nitrogen fertilizers are prohibited as are pesticides (except for fungicides).
- (4) Land group 4 (80–100 sinkholes/100 ha). Only meadows and forests are allowed. All fertilizers and pesticides are prohibited. In all the

land groups a 25 m radius protection zone is required around each doline. Within this protection zone, only grass without fertilizers or pesticides may be grown.

In addition to these measures, it is illegal to apply ammonium water and liquid ammonium to the soils of all four categories. It is also prohibited to use aircraft for spraying chemicals and mineral fertilizers. Ecologically sound agricultural plans have been designed for each land group. Biological agriculture is being introduced to the region. Thus, the protection of karst water from pollution and the reduction of human impacts on vulnerable karst groundwater is now official government policy. Funding from the national budget, therefore, is being provided to enable the implementation of the necessary protection measures (building of waste water treatment plants, manure storage facilities etc.). It is hoped that the introduction of these protection measures will stabilize karst development in the karst region of Lithuania.

The Tatula Board, named after the karst River Tatula, is officially responsible for the protection of the gypsum karst area and its important groundwater resources. It was established to comply with Resolution 589 of the Government of Lithuania, 24 December 1991. This resolution officially recognized the karst area and formalized protection and monitoring procedures in the area. The program of measures was adopted by the Decree of the Lithuanian Government by Resolution 719 on 17 September 1993. The Tatula Board is funded through the Ministry of Agriculture from central government funds. It encourages environmentally-friendly agriculture and antipollution measures in the gypsum karst area. It does this by organizing training courses at the local college, publishing advisory brochures and encouraging organic farming. It also tries to help the funding of water treatment plants for treating effluent. To limit the amounts of nitrate, phosphate and potassium entering the karst water, the Tatula Board helps farmers with interest-free loans for developing environmentally-friendly (organic) agriculture. The farmers have to produce a 5-year business plan and agree not to use insecticides and fertilizers. There are currently 43 farms working to sustainable bio/organic agriculture. These farms

are all in the third and fourth karst land groups with 50–80, or >80, sinkholes per 100 ha (km^{-2}). In all four agricultural categories, the law requires a 25 m zone of exclusion to agriculture around each sinkhole, and around some it requires an earth barrier to prevent runoff from entering the hole. The organic farming is monitored by the Society for Bio-Organic Agriculture (GAJA), which checks to see that no fertilizers are used. The Tatula Board has a program which uses 11 institutes and organizations to monitor environmental aspects such as groundwater. They would like a more extensive report to include waste water inspection and pollution control.

In contrast, to the protection of the karst water in Lithuania, some unsuitable practices have been noted in England. These include the piping of surface water run off from roads into sinkholes to drain it away. In addition, during the 1970s, some large sinkholes were filled with domestic refuse. Any leachate from this will have drained directly into the gypsum karst water system and may threaten local springs and wells supplying farms.

5. Conclusions

The presence of gypsum constitutes a geological hazard that can be mitigated by careful planning. This planning can be undertaken on several fronts and at several different scales. From a national point of view, the starting point is the recognition that the gypsiferous areas are special and potentially dangerous. At the local scale, the hazards are best considered as part of locally-applied planning and development processes. Local development plans should consider the implications of unstable land associated with gypsum karst and avoid the most unstable areas within and adjacent to sinkholes. This is the most cost-effective way of mitigating the hazard. Also on a local to regional scale the interaction of groundwater drawdown and recharge within the gypsum karst area needs to be considered. Since water extraction can trigger the subsidence, it should be carefully controlled.

Once the development areas have been defined, the hazard can be further mitigated by careful control of building and construction designs. The

use of special reinforced and extended foundations can be specified, along with protection to cables and pipes servicing the constructions. The implementation of these measures can be made effective by giving control to local authorities and requiring verification of the investigation and design procedures which are adopted in the subsidence-prone areas.

Where the gypsum karst is also closely related to the local potable groundwater supply, measures to protect the aquifer can be very beneficial. These measures can include limits on the types of agriculture and education or regulation to prevent pollution of the groundwater, especially through runoff or illegal drainage into sinkholes.

Acknowledgment

This study has been funded and facilitated by the Overseas Development Administration (ODA) of the British Foreign Office under Technical Development and Research (TDR) contract R6490-Gypsum Geohazards: their impact on development. Anthony H. Cooper particularly wishes to thank Drs. G. Motuza, J. Satkunas and K. Kadunas, plus all the staff at the Geological Survey of Lithuania for their help and hospitality. He wishes to give special thanks to Dr V. Marcinkevičius for explaining about the gypsum karst in Lithuania. Drs A.J. Reedman and J.D. Bennett are thanked for helping to arrange this project. T.J. Charsley and Dr J.H. Powell are thanked for reviewing the manuscript. Anthony H. Cooper publishes with permission of the Director, British Geological Survey (N.E.R.C.).

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